The Effect of Dry Unit Weight in Hydrodynamic Dispersion of Clayey Soils

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Abstract: In this study, hydrodynamic dispersion is explained which is the combined effect of grain diffusion and dispersion of a soil. Dry unit weight of a soil is shown to be the most effective parameter in hydrodynamic dispersion. The results obtained by traditional consolidation tests are compared to the theoretically obtained counterparts. All of the aforementioned theoretical values are derived from the theory of hydrodynamic dispersion. Additionally, by considering the hydrodynamic dispersion of a soil during consolidation, time rates of pore water, settlements and compression of soil skeleton can also be considered in terms of hydrodynamic dispersion and so the viscous properties of soils are able to be investigated as well. It can also be pointed out that the obtained data can especially support the studies about slope stability problems of soils.

Keywords: hydrodynamic dispersion, soil skeleton, time rates of settlement, dry unit weight, pore water, grain diffusion, dispersion

1. Introduction

Hydrodynamic dispersion is a phenomenon in which soil particles form a more disperse structure by means of orientation and soil grains diffuse because of very low time rates of settlement[1].

This phenomenon takes place in micro pores and at the outside of a solid-liquid intersection as a result of low time rates of pore water [2]. Variations in time rates of settlement are in a correlation with the changes in pore water pressures [3].

2. Material and Methodology

Soil samples which were taken from Mersin City in Turkey were used to perform a series of traditional consolidation tests. Since dry unit weight is the most effective parameter in hydrodynamic dispersion, soil samples with different initial and final effective stress variation ratios are presented in this study.

Results of these experiments are compared to theoretical values and the effect of dry unit weight in a consolidation is expressed [4].

The phenomenon was studied statistically at first and by performing a series of regression analyses, the most compatible equations were constructed.

Consequently, Equation 1 was given as the dispersion differential equation in terms of dry unit weight [2], [4]:

$$\frac{\partial \gamma_k}{\partial t} = v_z \frac{\partial \gamma_k}{\partial z} + D_s \frac{\partial^2 \gamma_k}{\partial z^2}$$
(1)

where, γ_k is dry unit weight, v_z is time rate of settlement, D_s is diffusivity coefficient. Coordinate z and time t are the

independent variables in Equation 1.

3. Results and Tables

Plasticity indices, dry unit weight variation ratios before and after a consolidation test, pre-consolidation pressures and over consolidation ratios of the two kind of clayey soil samples which were used in the performed tests are given below:

Soil Type 1 \Rightarrow $I_p = \%13$; $\Delta \gamma_k = \%10.98$; $\sigma_c = 2.3 \text{ kg/cm}^2$; OCR = 4.8 Soil Type 2 \Rightarrow $I_p = \%37$; $\Delta \gamma_k = \%7.78$; $\sigma_c = 2.2 \text{ kg/cm}^2$; OCR = 4.7

Results of the performed consolidation tests and diffusive and dispersive parameters of the Soil Type 1 samples can be seen in Table 1 and Table 2, respectively.

Results of the performed consolidation tests and diffusive and dispersive parameters of the Soil Type 2 samples can be seen in Table 3 and Table 4, respectively.

 Table 1: Effective stress variations relative to dry unit weights for the soil type 1 sample

Total Press. p (kg/cm²)	Samp. Height H (mm)	Void Ratio e (%)	Pressure Increment Δp (kg/cm ²)	Dry Unit Weight Yk (gr/cm ³)	Coeff. Of Volume Comp. m _v (cm ² /kg)	Efective Stress Increment σ' (kg/cm ²)
0.00	20.00	0.667	-	1.484	-	-
0.51	19.62	0.635	0.51	1.513	0.0373	0.52
1.02	19.33	0.611	0.51	1.536	0.0290	0.52
2.04	18.95	0.579	1.02	1.567	0.0193	1.04
4.08	18.39	0.533	2.04	1.614	0.0145	2.04
8.15	17.67	0.473	4.08	1.680	0.0096	4.17
4.08	17.80	0.483	4.08	1.668	0.0017	4.22
1.02	18.02	0.502	3.06	1.647	0.0041	3.09

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Total Press. p (kg/cm ²)	Dry Unit Weight γ _k (gr/cm ³)	Dif. Coeff. D₅×10 ⁻⁵ (cm²/dk)	Disp. Vrb. x	Disp. Flux Js (gr/cm²dk)	Dispers. Soil Amount ΔWs (gr)	Total Comp. Soil ΔWt (gr)
0.00	1.484	_	-	-	-	-
0.51	1.513	1.282	7.220	3.54×10 ⁻²⁹	1.0×10 ⁻²⁴	0.0216
1.02	1.536	2.210	5.418	6.48×10 ⁻¹⁹	1.8×10 ⁻¹⁴	0.0684
2.04	1.567	3.362	4.306	6.34×10 ⁻¹⁴	1.8×10 ⁻⁹	0.1711
4.08	1.614	4.928	3.452	9.06×10 ⁻¹¹	2.6×10 ⁻⁶	0.4109
8.15	1.680	6.714	2.841	7.46×10 ⁻⁹	2.1×10 ⁻⁴	0.8965
4.08	1.668	6.410	2.929	4.12×10 ⁻⁹	1.2×10 ⁻⁴	0.7946
1.02	1.647	5.877	3.097	1.27×10 ⁻⁹	0.4×10 ⁻⁴	0.6335

Table 2: Dispersion results of the soil type 1 sample

Note: Since consolidation is concluded at the end of each pressure increment, time rate of settlement v_z was taken as $v_z=0$.

 Table 3: Effective stress variations relative to dry unit weights for the soil type 2 sample

Total Press.	Samp. Height	Void Ratio	Pressure Increment	Dry Unit Weight	Coeff. Of Volume Comp.	Efective Stress Increment
р	н	e	Δр	Yk	mv	σ
(kg/cm²)	(mm)	(%)	(kg/cm²)	(gr/cm ³)	(cm²/kg)	(kg/cm²)
0.00	20.00	0.429	-	1.710	-	-
0.51	19.85	0.418	0.51	1.724	0.0146	0.56
1.02	19.62	0.401	0.51	1.744	0.0228	0.51
2.04	19.24	0.374	1.02	1.779	0.0190	1.05
4.08	18.71	0.336	2.04	1.829	0.0136	2.04
8.15	18.00	0.286	4.08	1.901	0.0093	4.15
4.08	18.13	0.295	4.08	1.887	0.0017	4.35
1.02	18.57	0.326	3.06	1.843	0.0080	2.95

Table 4: Dispersion results of the soil type 2 sample

Total Press.	Dry Unit Weight	Dif. Coeff.	Disp. Vrb.	Disp. Flux	Dispersive Soil Amount	Total Compacted Soil
p (kg/cm²)	γk (gr/cm³)	D₅×10⁻⁵ (cm²/dk)	x -	J₅ (gr/cm²dk)	∆W s (gr)	Δwv _t (gr)
0.00	1.710	-	-	-	-	-
0.51	1.724	0.515	11.525	9.75×10 ⁻⁶⁵	2.8×10 ⁻⁶⁰	0.0041
1.02	1.744	1.282	7.220	4.16×10 ⁻²⁹	1.2×10 ⁻²⁴	0.0254
2.04	1.779	2.490	5.080	3.17×10 ⁻¹⁷	9.0×10 ⁻¹³	0.1029
4.08	1.829	4.052	3.873	3.44×10 ⁻¹²	9.7×10 ⁻⁸	0.3013
8.15	1.901	5.927	3.081	1.65×10 ⁻⁹	4.7×10 ⁻⁵	0.7499
4.08	1.887	5.602	3.192	0.74×10 ⁻⁹	2.1×10 ⁻⁵	0.6497
1.02	1.843	4.441	3.672	0.02×10 ⁻⁹	0.6×10 ⁻⁶	0.3733

Note: Since consolidation is concluded at the end of each pressure increment, time rate of settlement v_z was taken as $v_z=0$.

Correlations between experimental and theoretical effective stress variations are given in Figure 1 and Figure 2 for the Soil Type 1 and Soil Type 2 samples, respectively.

The aforementioned theoretical values were obtained by using Equation 2 given below which was derived from the solutions of the differential equation given as Equation 1 [2], [4].

$$\sigma' = \frac{1}{m_v} \ln \frac{\gamma_{k2}}{\gamma_{k1}} \tag{2}$$

where, m_v is the coefficient of volume compressibility, γ_{k1} and γ_{k2} are the initial and final dry unit weights, respectively.

It is clear from the correlations in Figure 1 and Figure 2 that theoretically obtained values are very close to the experimental counterparts.







Figure 2: Correlation of theoretical ($\Delta \sigma$ ') and experimental (p) effective stress increments for the soil type 2 sample

The comparison between diffusivity coefficients (D_s) of the Soil Type 1 and Soil Type 2 samples are seen in Figure 3 below.



Figure 3: Correlation of diffusivity coefficients for the soil type 1 and soil type 2 samples

As a result of a higher ratio of dry unit weight variation of the Soil Type 1 sample throughout the performed tests,

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diffusivity coefficients for the Soil Type 1 samples are found to be higher than that for the Soil Type 2 samples.

The obtained values for diffusivity coefficient were obtained by using Equation 3 which was derived from the solutions of the differential equation given as Equation 1 [4].

$$D_s = \frac{z^2}{4t} \ln \frac{z_i}{z} \tag{3}$$

where, z_i ve z are the initial and any sample heights, respectively, $z^2/4$ is the squared power of the drainage path z/2 and t is time [4].

The values for flux (J_s) and dispersive variable (x) were found by using Equation 4.a and Equation 4.b, respectivelywhich were also derived from the solutions of the differential equation given as Equation 1, as well[4]:

$$J_{s} = \sqrt{\frac{D_{s}}{\pi t}} \cdot \left(\gamma_{kf} - \gamma_{ki}\right) e^{-x^{2}}$$
(4.a)

$$x = \frac{z + v_z t}{2\sqrt{D_s t}} \tag{4.b}$$

where, D_s is the diffusivity coefficient, γ_{ki} and γ_{kf} are the initial and final values for dry unit weights, respectively, v_z is the time rate of settlement, z is the sample height at any instant and t is time[4].

The values for the amount of soil filling into macro pores and diffusing inside the soil matrixwere found by Equation 5 below[4]:

$$\Delta W_{\rm s} = J_{\rm s} A \Delta t \tag{5}$$

where, J_s is flux, A is the cross sectional area and t is time[4]. The correlation between total compacted soil amounts(ΔW_t) of the Soil Type 1 and Soil Type 2 samples can be seenin Figure 4 below.



Figure 4: Total compacted soil amounts (Δw_t) for the soil type 1 and soil type 2 samples

The values for the total compacted soil amounts were obtained by Equation 6 below[4]:

$$\Delta W_t = A \Delta z \Delta \gamma_k \tag{6}$$

where, A is cross sectional area, z is deformational height and γ_k is dry unit weight[4].

The equations for obtaining the values for dispersed (ΔW_s) and total (ΔW_t) soil amounts which are given as Equation 5 and Equation 6, respectively were also derived from the solutions of the differential equation given in equation 1 [2], [4].

If the correlation given in Figure 4 is observed, it is seen that all of the values for the Soil Type 1 samples are higher than that for the Soil Type 2 samples.

This relation is also the cause of the effect of dry unit weight variation percentages which is higher for the Soil Type 1 samples.

4. Conclusion

This study presents information on the subject of hydrodynamic dispersion and the effect of dry unit weight variations on the consolidation of clayey soils.

The equations for finding variations in effective stresses, diffusivity coefficients, dispersive flux, dispersive and total compacted soil amounts and if deserved time rates of settlements, time rates of pore water and time rates of compression of soil skeletoncan be obtained as a result of solving the differential equation given as Equation 1.

The aforementioned equations were shown to give results that are in a very close agreement to the experimental counterparts.

On the other hand, diffusive and dispersive characteristic of a specific region in Mersin Turkey was investigated in this study and the effects of dry unit weight were discussed in order to clarify hydrodynamic dispersion in a consolidation phenomenon.

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Author Profile



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